

A historical note on the outflows of the Okavango Delta System

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Periodic
+ timing of outflows
+ use of
systems in
years "near"

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Introduction

As pressure on land and water resources in Botswana has increased, consideration has been given to the utilisation of the Okavango Delta region. The use of such a complex and delicate ecosystem requires careful planning, a fact recognised by the UNDP/FAO (1977) study of the Okavango as a water resource.

One project arising from the UNDP/FAO study has been the formulation of a mathematical model to quantify and predict changes in the hydrological network (Dincer et al. 1976). This, of necessity, is limited by the spatial and temporal limitations of available accurate data; in this case to the period since 1952.

As the delta system is constantly in a state of adjustment to natural and man-made changes the value of the mathematical model can be supplemented by consideration of the historical record. This paper summarises what is known of the Okavango Delta outflows for the period 1849-1983, drawn from the observations of travellers, administrators and engineers, together with available aerial photography (since 1937) and satellite imagery (since 1972).

The principal sources used in this paper are listed in the bibliography. It is not possible to quote the source of every observation, nor has the list of potential sources been exhausted.

The data used is non-quantitative and varies in reliability. Remote sensing data and the direct observations of experts are reliable, whilst eye-witness recollections vary in quality. For this reason the earlier part of the record (particularly before 1925) must be interpreted with caution.

Physical Background

The Okavango Delta is shown in figure 1, whilst the outflows from the system are indicated for the historical period in figure 2. The Okavango Delta covers an area of approximately 22,000 km², receiving an average annual inflow of 10.5×10^9 m³ (UNDP/FAO vol 1:1). About 95% of this water is lost within the Delta by evapo-transpiration or groundwater recharge, the remainder drains towards the Zambezi River or one of the three lake basins (Lake Ngami, the Mababe

Depression and the Makgadikgadi Pans) which form the end-sinks of the endoreic drainage system. The amount of outflow varies from year to year with change in inflow, precipitation and hydrological condition in the swamps. The direction of outflow may be further influenced by changes within the swamps due to blockage (natural, animal or man-made) or local tectonic events.

A useful summary of the hydrology of the Okavango is given by Wilson & Dincer (1977) whilst a discussion of Okavango inflows can be found in UNDP/FAO vol 1:62-68.

Extensive lake shorelines and abandoned swamps indicate regime changes during the Quaternary (Cook 1980), but these are regarded as beyond the scope of the present inquiry.

The Outflows

Considered in an anti-clockwise direction the outflow are as follows:

(1) The Magwegqana (Selinda) Spillway connects the Okavango to the Linyanti (and hence to the Zambezi) along a channel some 120 km in length at a gradient of 1:4,000. The channel was not visited by Europeans until late in the 19th Century due to its inaccessibility. Early accounts gave conflicting opinions on the direction of flow. Landsat imagery clearly shows the bulk of the water coming from the Okavango, meeting backflow from the Linyanti about 10 km upstream of the Linyanti confluence.

The Spillway floods sporadically in "good" years (a ratio of 1:3 from the limited data available). There is no evidence to suggest that any changes in the regime have occurred in the historical period.

(2) The Savuti channel connects the Linyanti to the Mababe Depression, a distance of 94 km at a gradient of 1:5,200. Strictly speaking it is not an outflow of the Okavango under its present regime, although Du Toit (1925:19) noted a connection between the Magwegqana and Savuti channels.

The Savuti appears to have flowed frequently until the late 1880's, then dried up until 1958, with flow in most years subsequent to this. Stigand (1923:415) has commented on the synchronous desiccation of the Savuti and Thaoe systems, but they may have different causes. Certainly the absence of water in the Savuti

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channel during the high floods of 1925 suggests that tectonic rather than hydrologic factors are involved.

The absence of a substantial body of water in the Mababe Depression during the period of European expansion suggests that Savuti flows 1849–1890 were no greater than in recent years. This may not have been the case in the 18th Century, when oral tradition supports the existence of a "Lake Mababe" (Flou 1972:152).

(3) The Mababe River conveys water from the Nqogha branch of the Okavango to the southern extremity of the Mababe Depression. In fact the water arrives by two routes, the Moanachira-Khwai-Mochaba-Kudumane system flowing to the east, and the Gomoti-Mghelo-Zamkuio system flowing in an arc further south. The two rivers join 8 km from Mababe, but have not been differentiated in figure 2.

The Mababe fed a substantial swamp in the 19th Century, and oral reports (Campbell & Child 1971) describe a river deep enough to hinder passage. The river dried up around 1910, and has subsequently reached the Depression only in 1925 and 1957, although water was present at the Kudumane-Zamkuio confluence (upper Mababe) on several occasions in the 1950s.

(4) The Thamalakane system is the major outflow of the Okavango, collecting water from the central distributaries along the Thamalakane Fault, then bifurcating below Maun into the Lake (Nghabe) and Boteti Rivers, flowing to Lake Ngami and the Makgadikgadi Pans respectively. The ratio of flow between the two rivers was 2:1 in favour of the Boteti until the construction of a bund 2 km below the bifurcation in 1969; present flow ratios are about 5:1 (Kraatz 1976). As the annual flood reaches the upper section of both rivers every year (with the possible exceptions of 1933 and 1935–36) Toteng and Rakops have been used as gauging points to reveal more subtle variations in flow.

(5) The Kunyere River flows along a fault line from the Xudum system to Lake Ngami, joining the Lake River at Toteng. The river reaches Toteng in most years, but carries a smaller volume of water than the Lake River in a "good" year.

(6) The Thaoge River entered Lake Ngami from the north until the early 1880's (Shaw 1983) although the flow in the system, together with the level of Lake Ngami, had been declining during and prior to the period under consideration (Shaw, in press). Unlike the other outlets the Thaoge was originally represented by a series of anastomosing channels (i.e. true swamp) rather than a single distributary, yet it is probable that only one outflow was operative by the 1850's. The gradual diminution of flow over decades suggests that natural and man-induced blockages were a major factor in this system.

Conclusion

The pattern of outflows shown in figure 2 can be broken down into 3 episodes:

(1) 1849–c.1900: The period starts with the function of all seven outlets at current average or above average flows, and ends with the drying up of Lake Ngami, the lower Thaoge, the Savuti, and, slightly later, the Mababe River.

(2) 1900–1951: This period coincides with the incidence of limited flows throughout the remainder of the system, interspersed with good years such as 1910, 1925–27 and 1944. The decade of the 1930's saw a very dry period in the history of the Delta.

(3) 1951 to Present: The last thirty years have seen good but cyclic flows in the Thamalakane system, a resumption of flow in the Savuti and the occasional presence of water in the Mababe River. The period ends at 1983 with the lowest inflow on record and outflow conditions equivalent to 1933.

The mathematical model developed by the UNDP/FAO Project and employed by the Department of Water Affairs is based within the third period, and uses selected portions of the Okavango Delta for which quantitative data is available. The model has shown the existence of regime changes at approximately decadal intervals. These changes are not predictable and are a limiting factor in long-term water-use planning (UNDP/FAO vol. 1:85). The historical record shows that long-term changes of greater magnitude occur outside the reference of the model, some of them non-cyclic within the given time span. These must certainly have important consequences for the utilisation of the Delta region.

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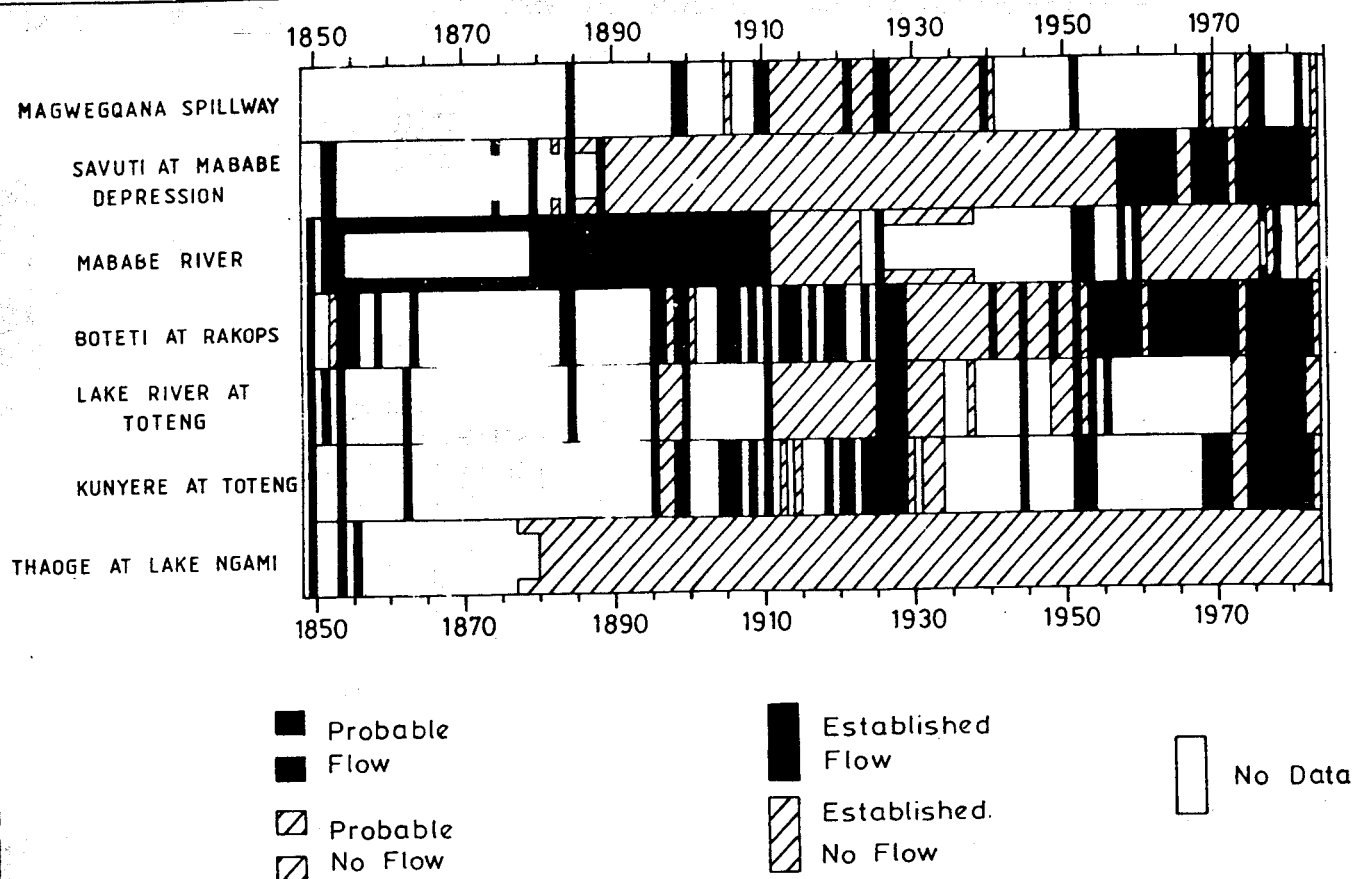


Figure 2 Outflows from the Okavango System 1849-1983

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